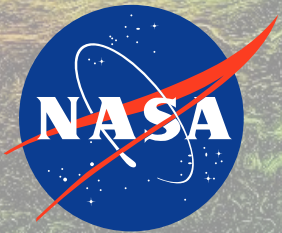
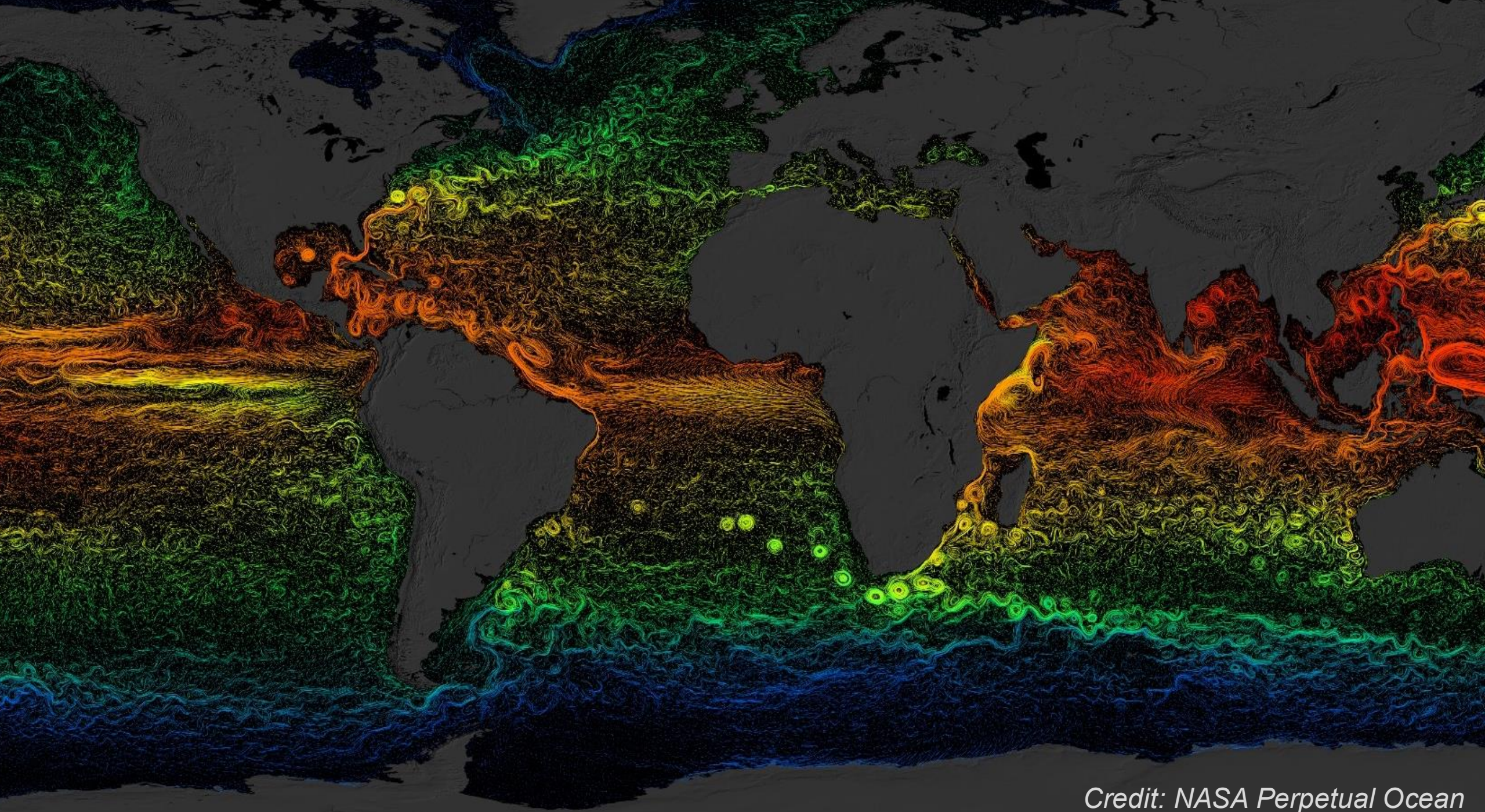


How the Ocean Redistributes Heat



Anthony Meza
MIT-WHOI Joint Program
AOML/PhOD Seminar

May 19 2026



Credit: NASA Perpetual Ocean

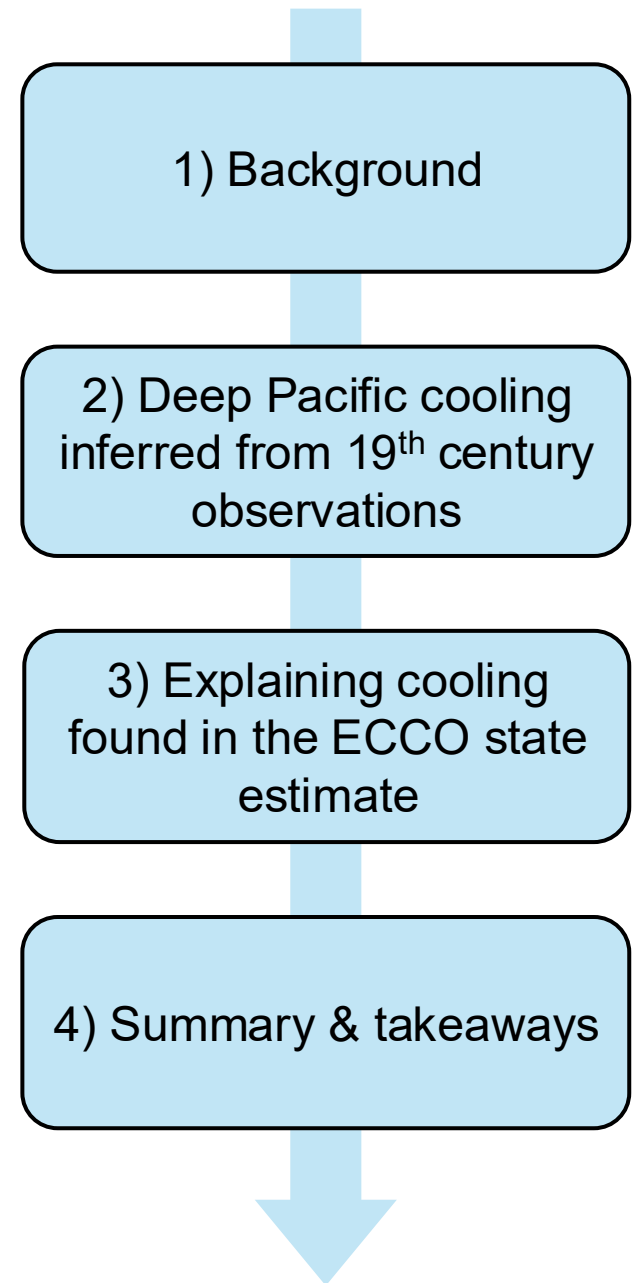
Introduction and outline



My Ph.D. thesis focuses on global ocean circulation variability and how it shapes abyssal tracer distributions.

My tools:

- High-res coupled climate models
- Ocean & atmospheric reanalyses
- Satellite & in-situ observations
- Idealized & statistical modeling



The ocean stores and redistributes heat

Seo et al., 2023

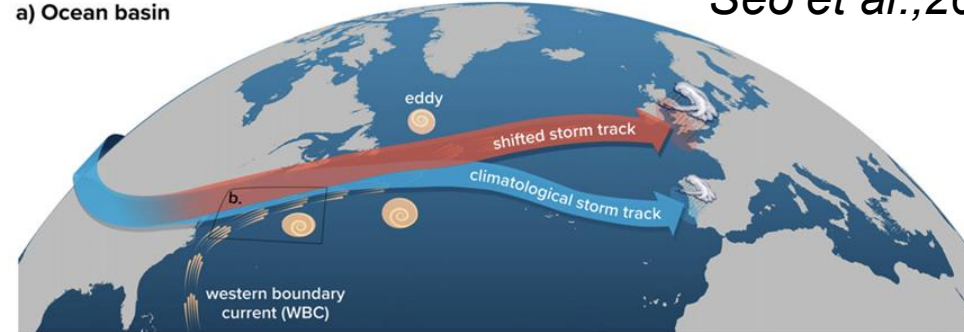
Heat enters and leaves at the surface

Large-scale currents redistribute heat in the interior

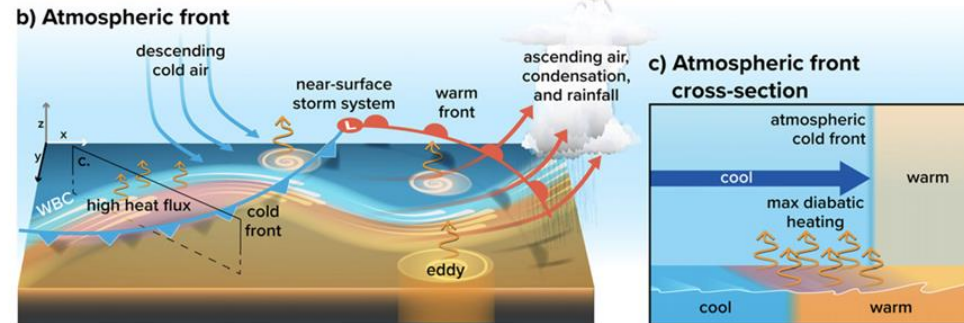
Eddies and turbulence mix heat across gradients

Ocean pathways influence heat re-emergence

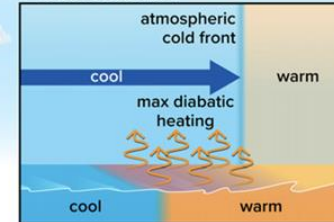
a) Ocean basin



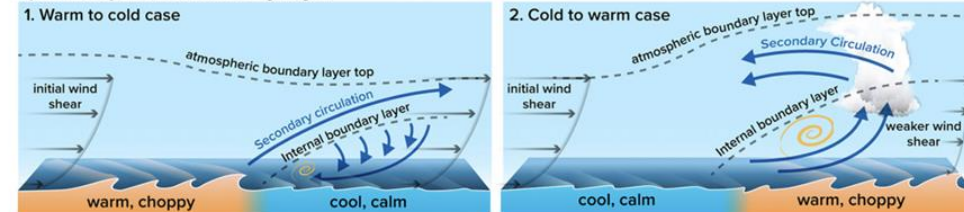
b) Atmospheric front



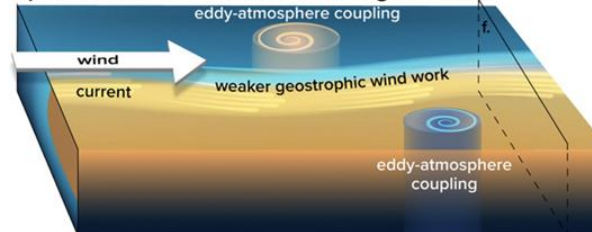
c) Atmospheric front cross-section



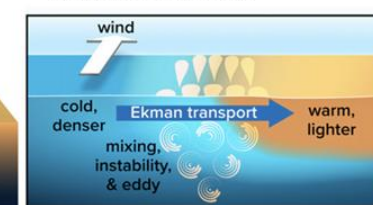
d) Atmospheric boundary layer



e) Ocean fronts & eddies interacting with wind



f) Stratification, instability & turbulence at fronts



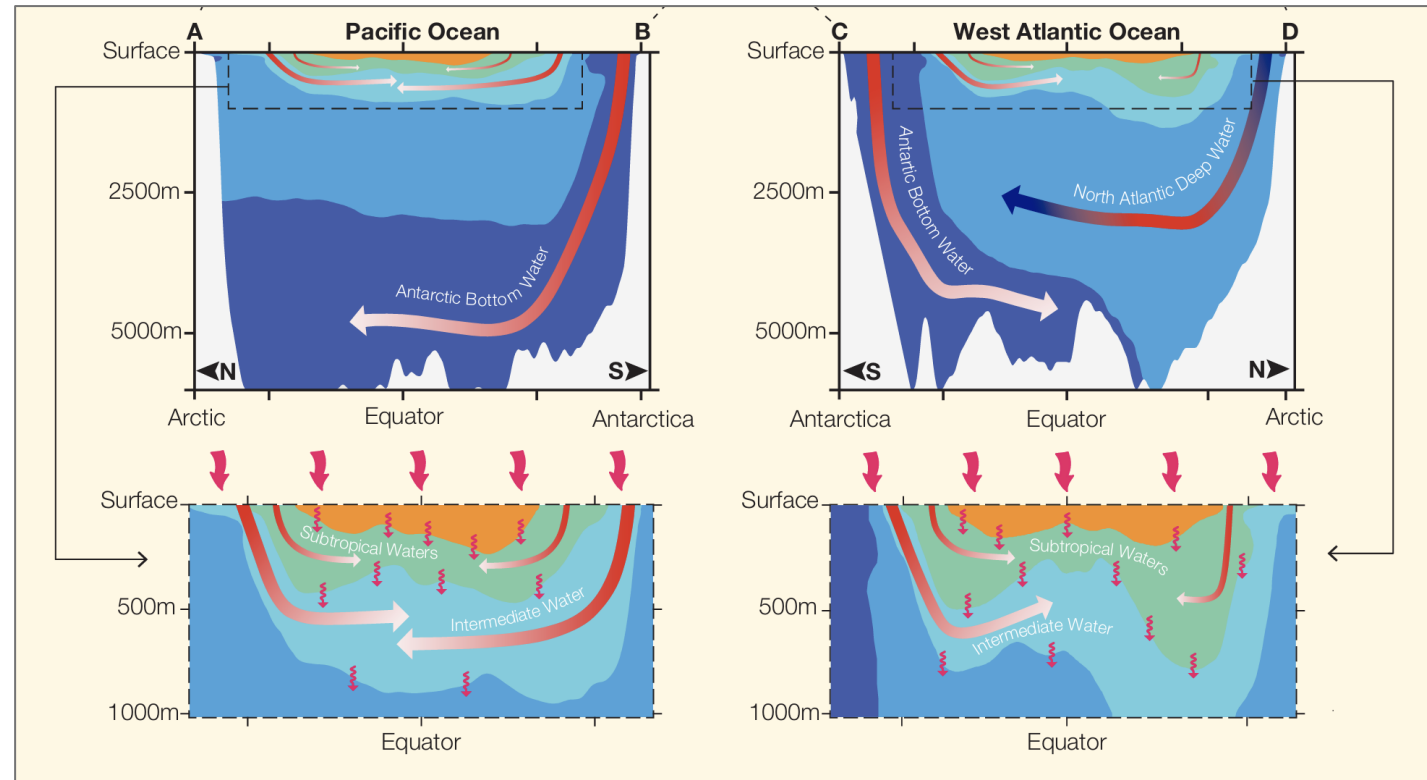
The ocean stores and redistributes heat

Heat enters and leaves at the surface

Large-scale currents redistribute heat in the interior

Eddies and turbulence mix heat across gradients

Ocean pathways influence heat re-emergence



IPCC AR5 FAQ 3.1, Figure 1

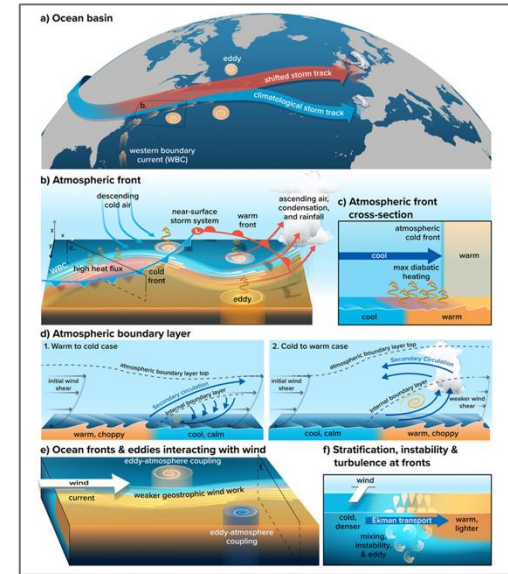
The ocean stores and redistributes heat

Heat enters and leaves at the surface

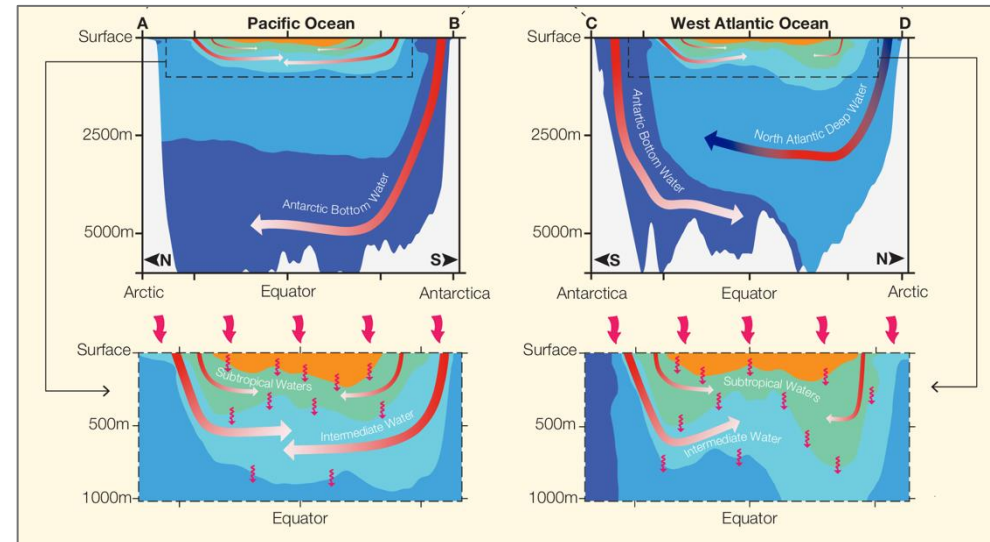
Large-scale currents redistribute heat in the interior

Eddies and turbulence mix heat across gradients

Ocean dynamics influence heat storage and re-emergence



+



The ocean stores and redistributes heat

Heat enters and leaves at the surface

Large-scale currents redistribute heat in the interior

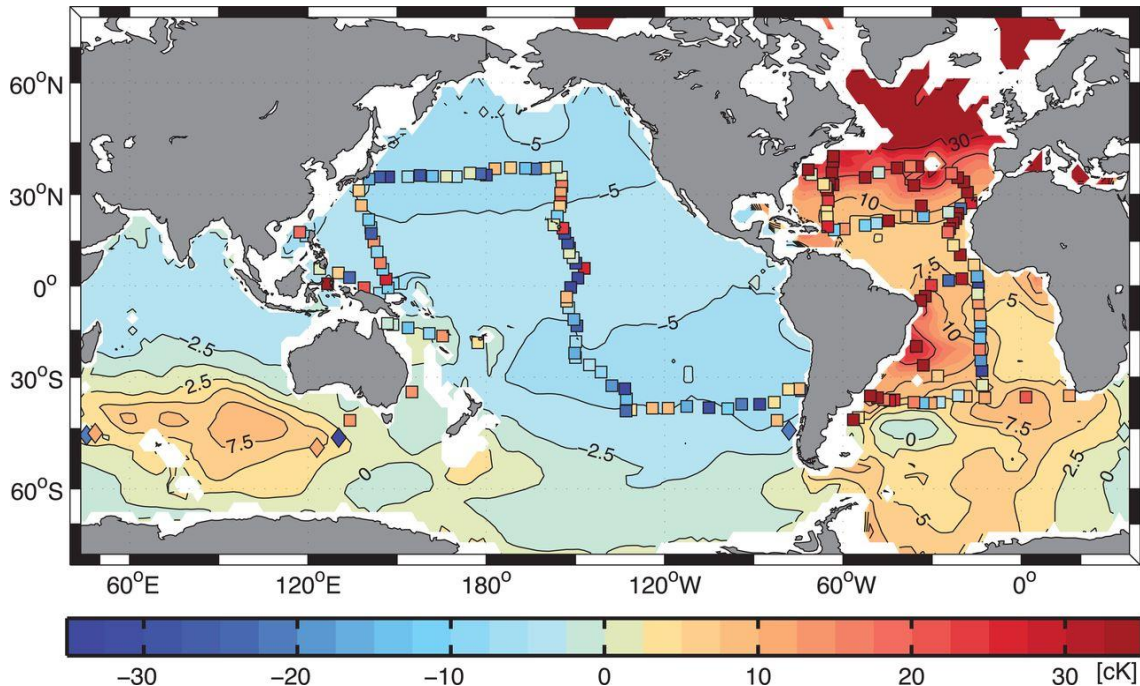
Eddies and turbulence mix heat across gradients

Ocean pathways influence heat re-emergence

**This
work**

A century of deep ocean temperature trends

Deep temperature differences (1870–1990)



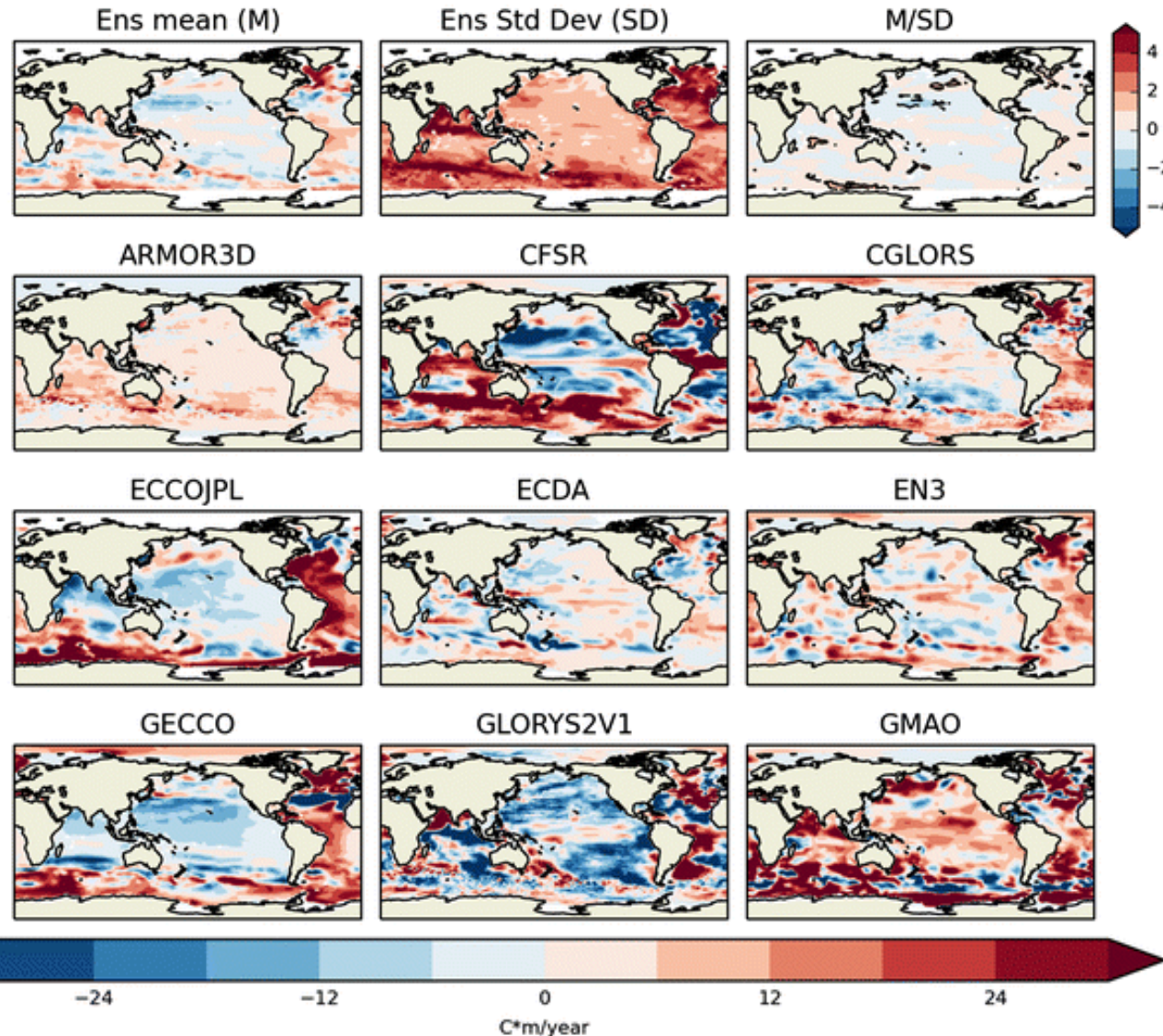
1 cK = 0.01 deg C

Gebbie & Huybers 2019

- 5 ± 3 cK/century cooling trend in deep Pacific inferred from HMS *Challenger* data
- A fixed-circulation inverse model explained the trend as a response to Little Ice Age surface temperature anomalies
- The inversion also suggested that the signal could still be in the modern ocean

Trend uncertainty among ocean reanalyses

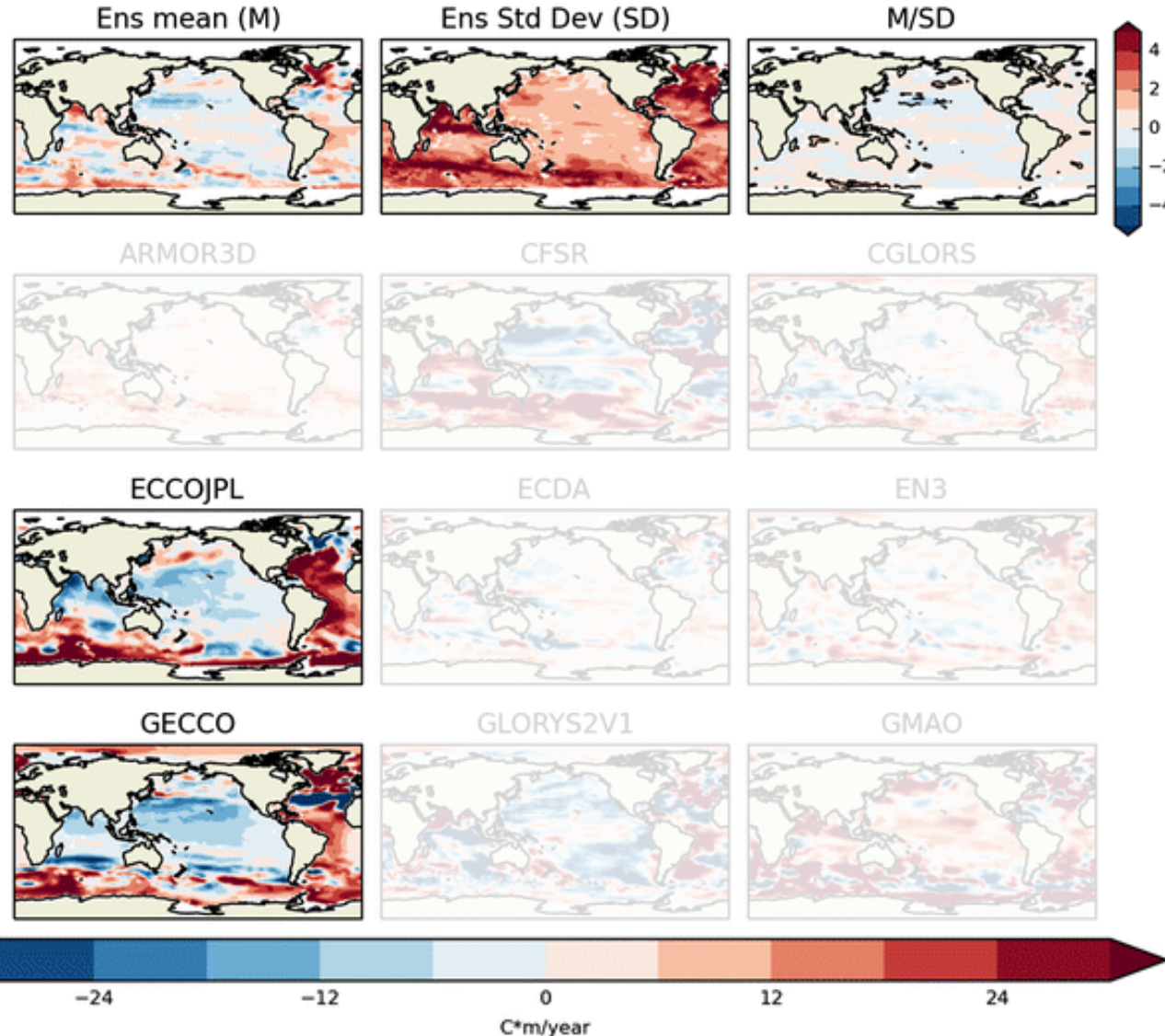
Trends in 700-6000m OHC (1997-2009)



*Adapted from
Palmer et al., 2017*

Trend uncertainty among ocean reanalyses

Trends in 700-6000m OHC (1997-2009)



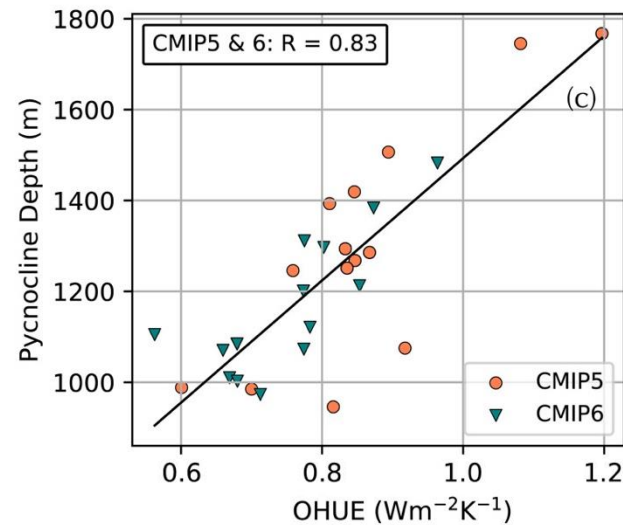
- Previous ECCO ocean state estimates show cooling in the deep Pacific

Adapted from Palmer et al., 2017

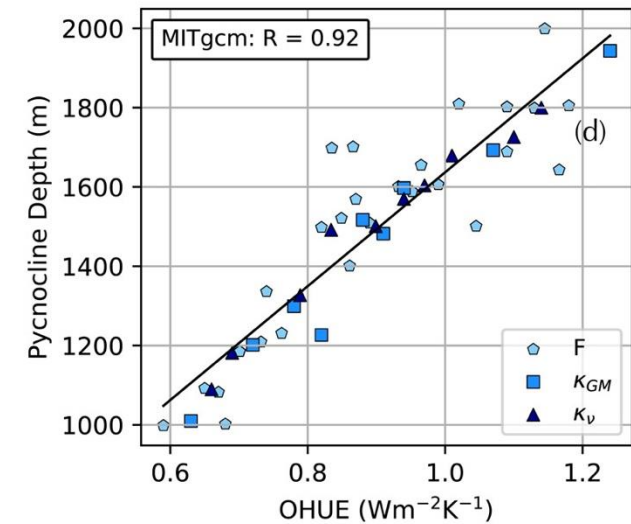
Initialization matters for climate simulations

- There is growing evidence that background temperature and salinity fields constrain ocean heat uptake

Multi-model comparison between pycnocline depth and ocean heat uptake



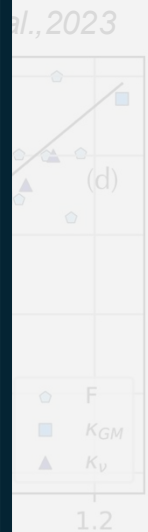
Newsom et al., 2023



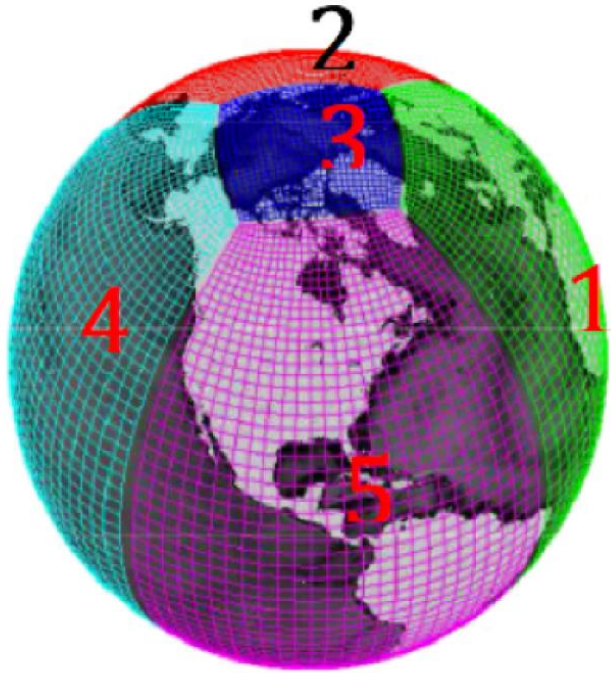
Initialization matters for climate simulations

Research questions:

- Does ECCO still have deep Pacific cooling?
- Does this signal emerge from constraining the model to observations?
- What physical mechanisms explain the cooling?



ECCO Version 4 release 4



ECCO Version 4 release 4 (V4r4)

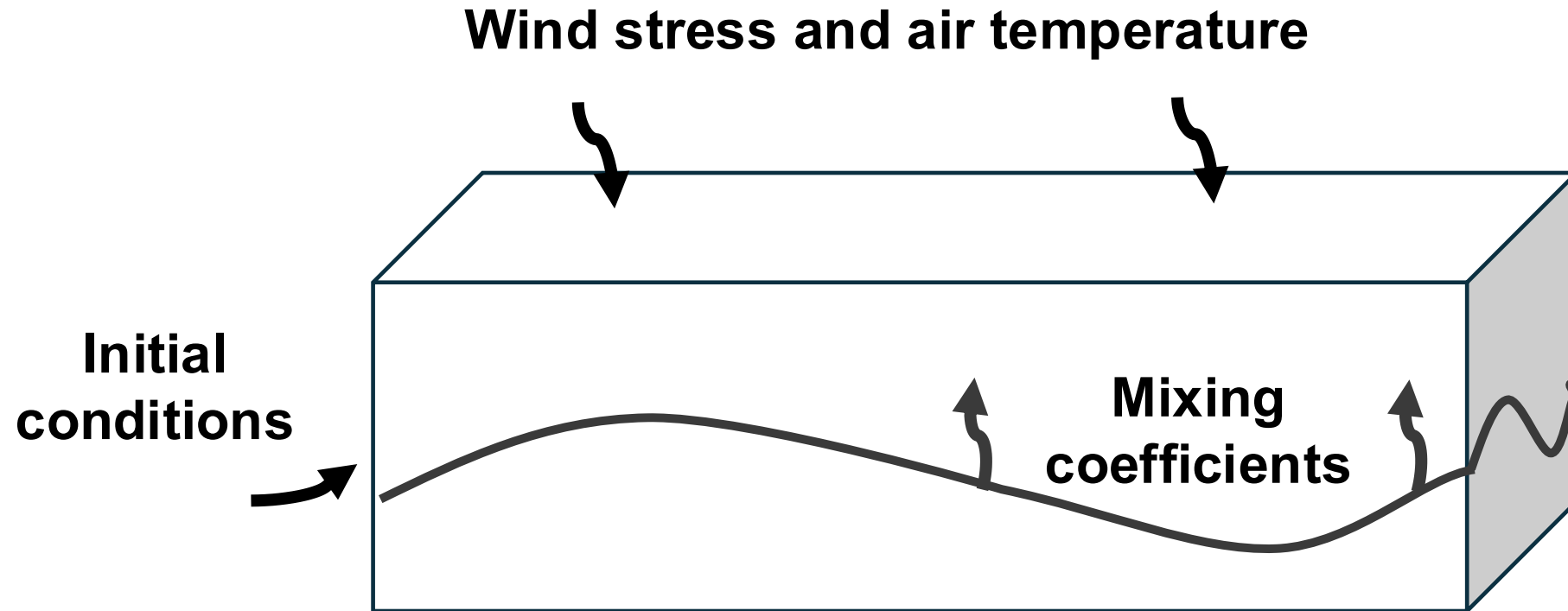
- Ocean state estimate spanning 1992–2017
- Nominal 1° resolution with 50 vertical levels
- Constrained by millions of observations

First-guess model configuration:

- Atmospheric forcing from reanalysis winds and air temperature
- Initial condition from the WOCE 2009 climatology
- Uniform mixing parameters



How ECCO V4r4 works



ECCO changes a few sets of parameters to fit to millions of ocean observations

Experimental design

- Initial set of perturbation experiments
- These quantify the effect of the adjustments used in Iteration 129 on temperature variability

Experiment Name	Control Adjustments Applied			
	Initial Condition	Wind Stress	Buoyancy Forcing	Mixing Parameters
Iteration 129 (<i>ECCO V4r4</i>)	✓	✓	✓	✓
Adjusted Initial Conditions	✓	×	×	×
Adjusted Wind Stress	×	✓	×	×
Adjusted Buoyancy Forcing	×	×	✓	×
Adjusted Mixing Parameters	×	×	×	✓
Iteration 0 (<i>ECCO V4r4 First Guess Solution</i>)	×	×	×	×

Experimental design

Mixing adjustment effect \approx Adjusted mixing parameters – Iteration 0

Experimental design

$$\text{Iteration 129} \approx \text{Iteration 0} + \sum \text{Effects}_i$$

Experimental design

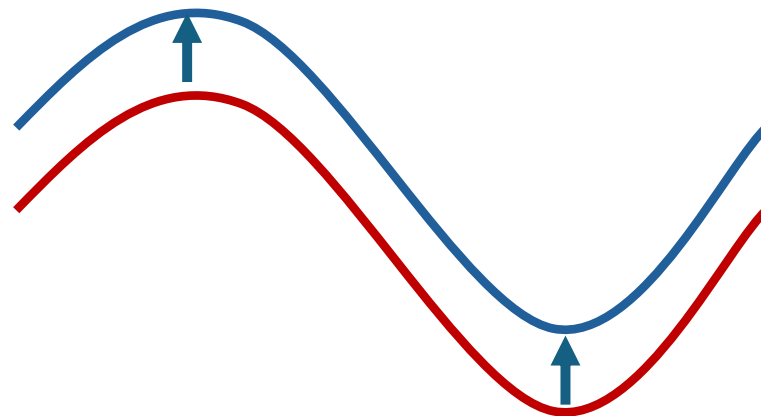
Experiment Name	Wind Stress Fields			
	First-Guess	First-Guess	Adjusted	Adjusted
	Seasonal Cycle	Extraseasonal	Seasonal Cycle	Extraseasonal
Seasonal Adjusted Wind Stress	×	×	✓	×
Seasonal First-Guess Wind Stress	✓	×	×	×

- These quantify the effects of Iteration 129 mean seasonal cycle adjustments on temperature variability

Experimental design

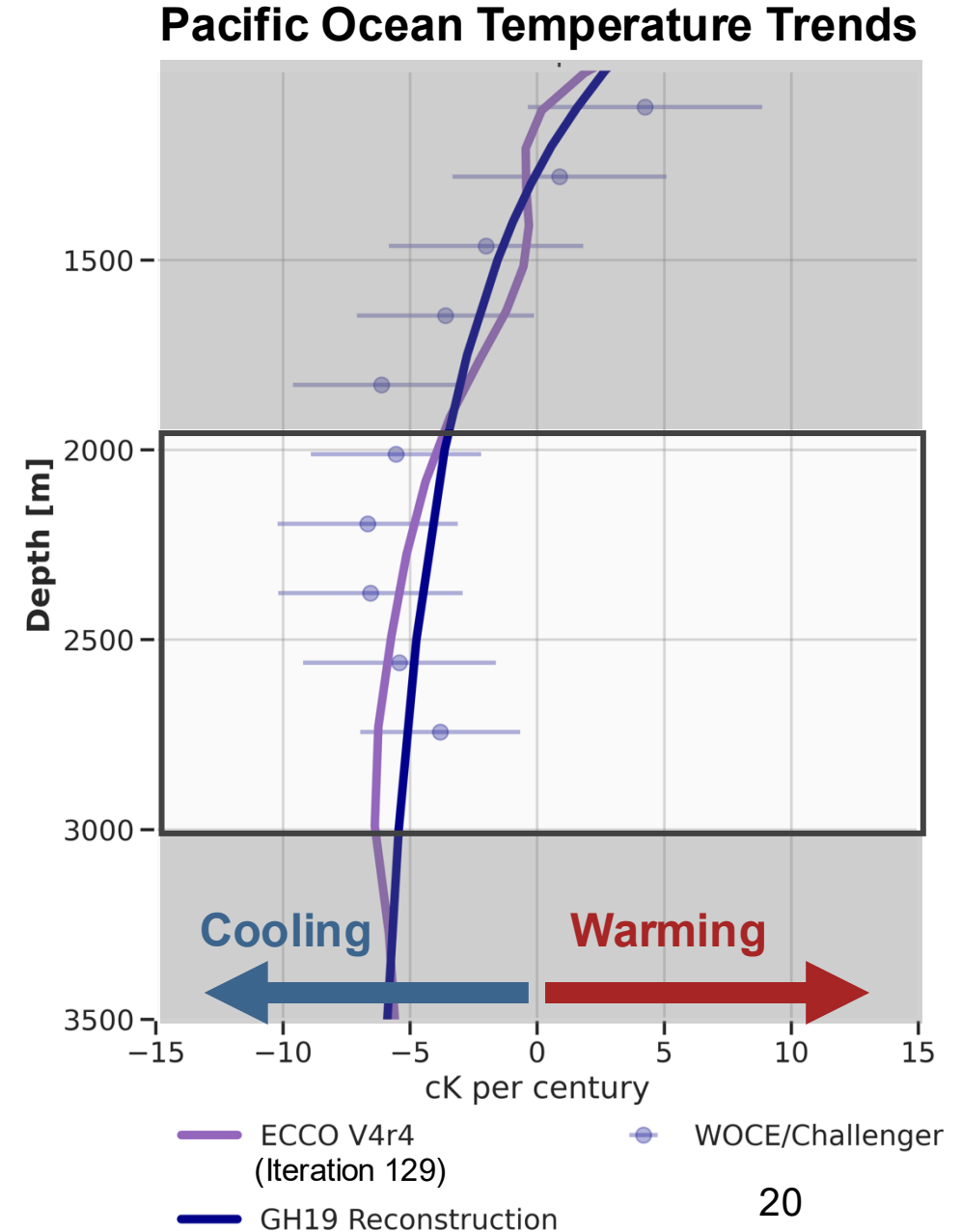
Mean seasonal cycle adjustment effect \approx Seasonal adjusted wind stress $-$ Seasonal first-guess wind stress

Example adjustments



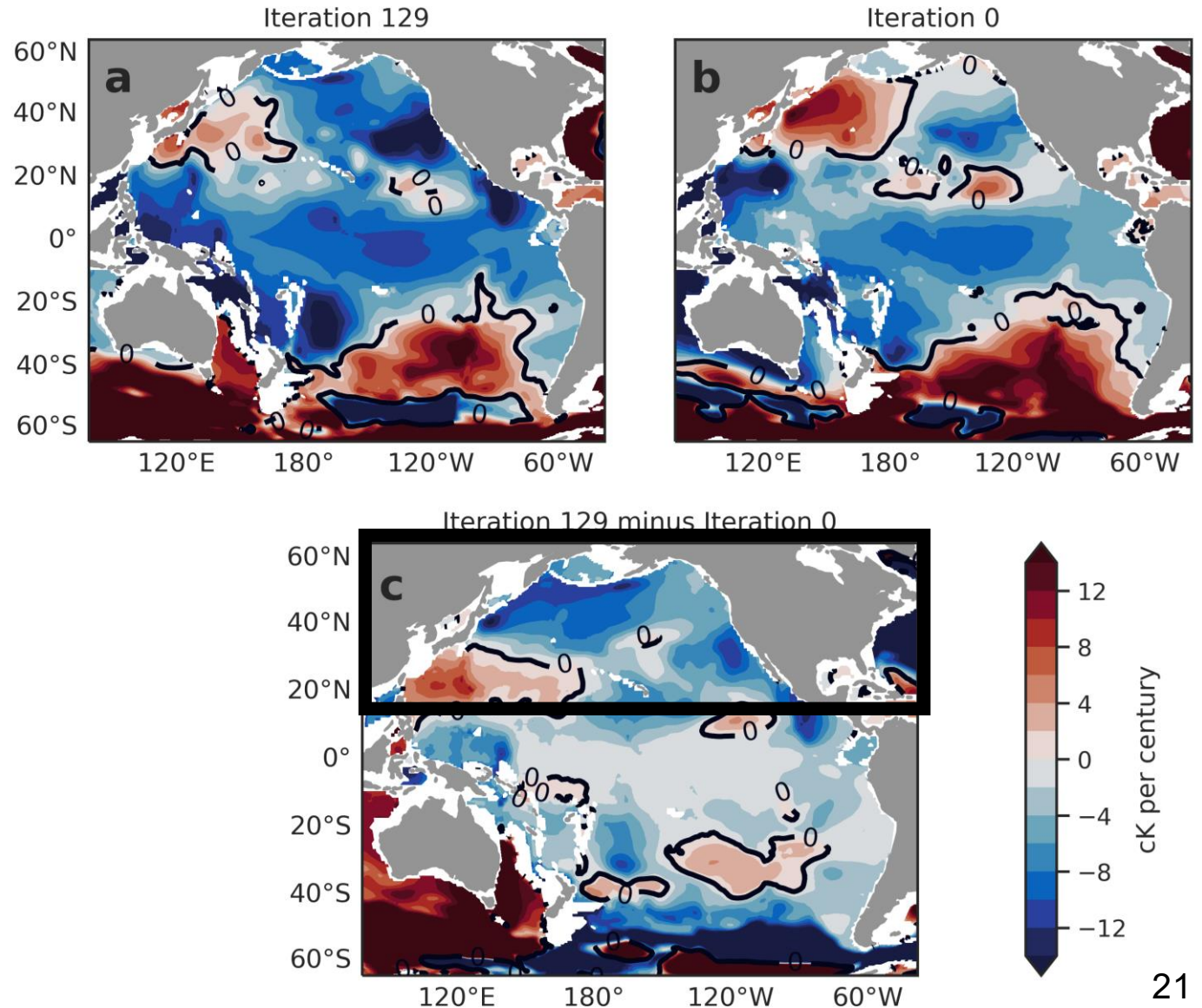
Pacific cooling in ECCO V4r4

- Below 1000 meter, ECCO cooling trends are comparable to the HMS *Challenger*/WOCE temperature differences



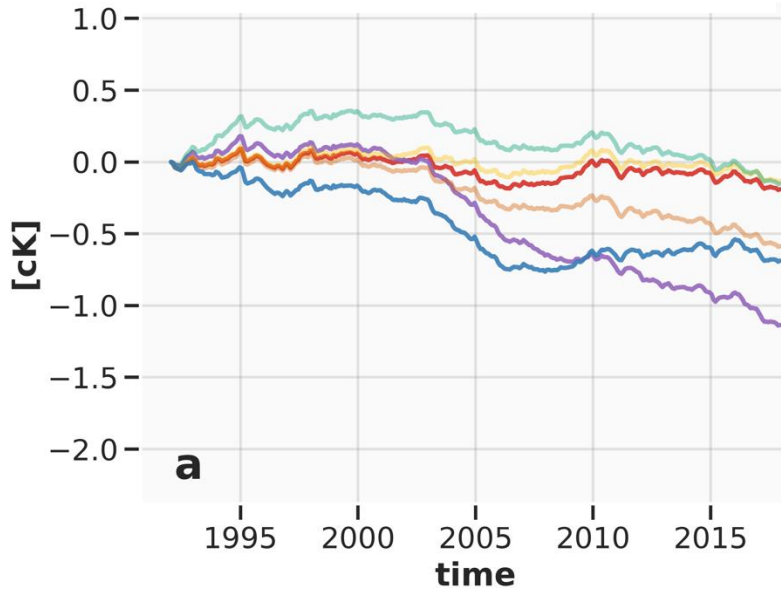
Mid-depth (2-3km) Pacific temperature trends

- Deep Pacific cooling intensifies after fitting ECCO V4r4 to data
- Enhanced North Pacific cooling is especially apparent



Mid-depth North Pacific heat budgets

Mid-depth North Pacific temperature anomaly

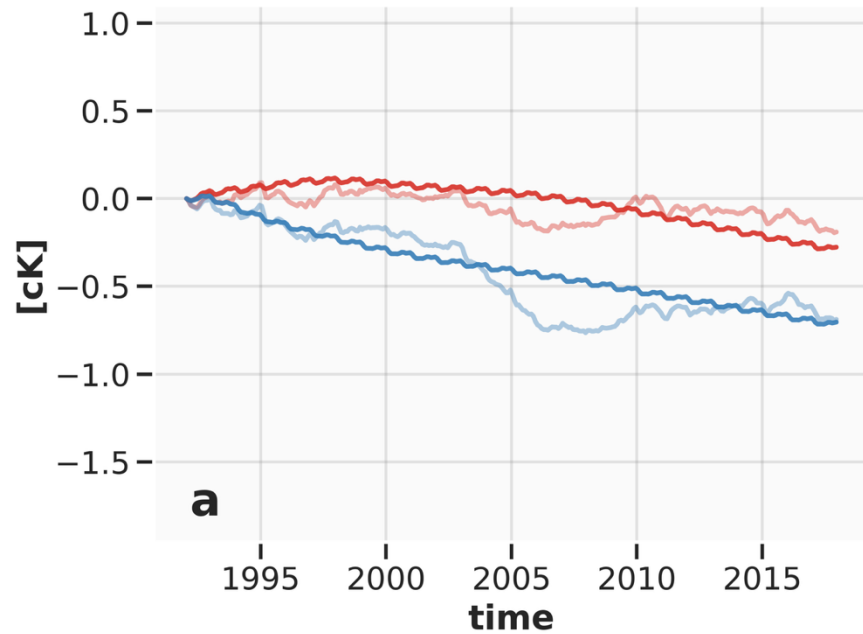


- Data-driven wind adjustments cause most of the cooling

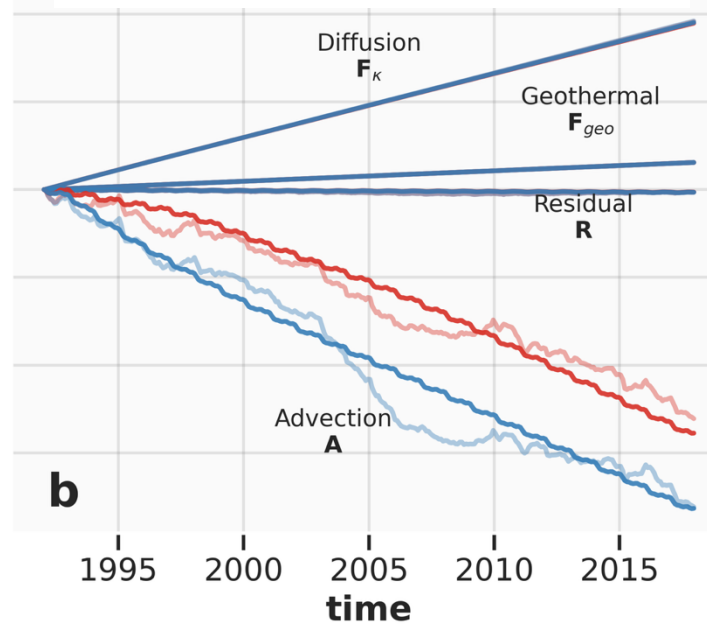


Mid-depth North Pacific heat budgets

Mid-depth North Pacific temperature anomaly



Heat budget

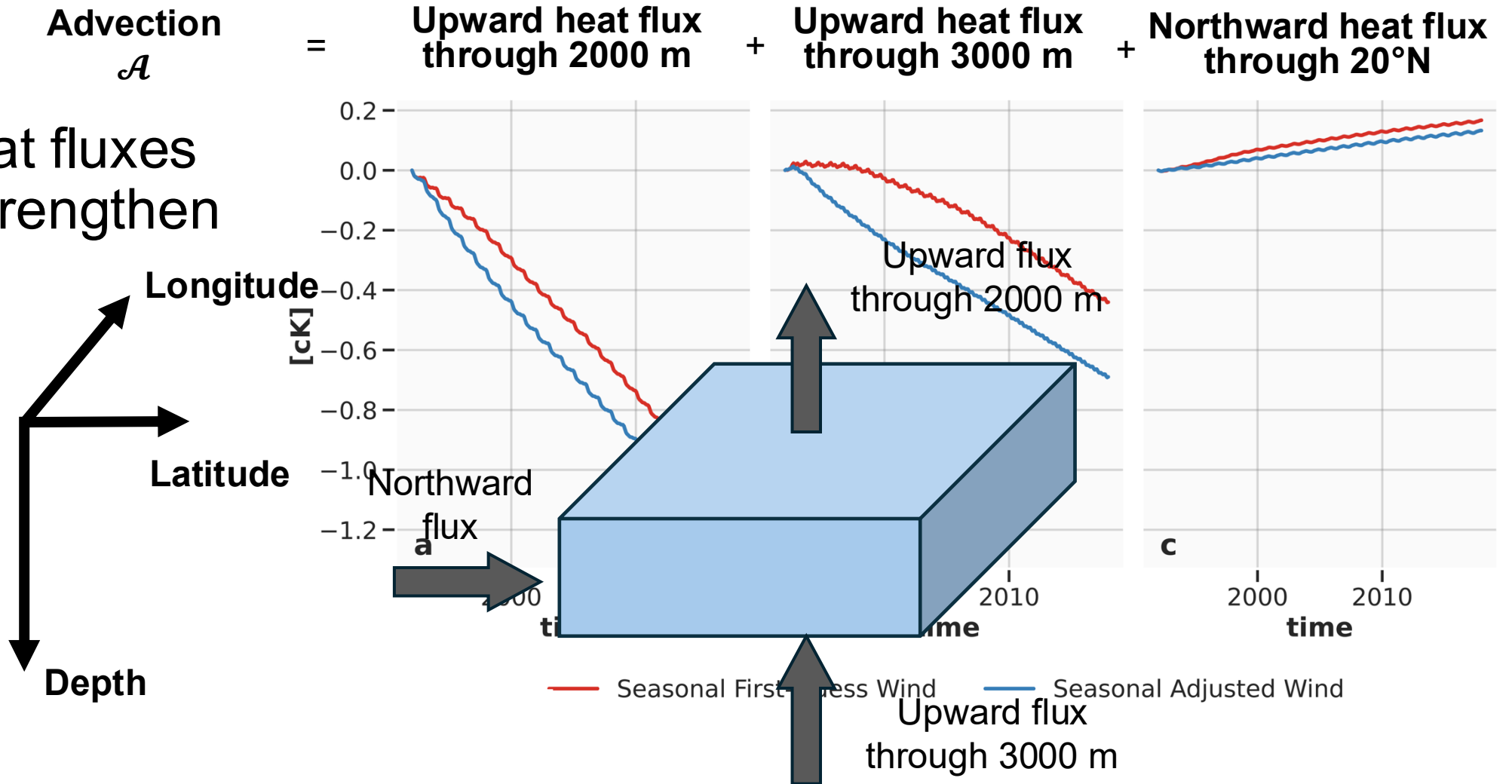


- Iteration 0 (ECCO V4r4 First-Guess)
- Adjusted Wind
- Seasonal First-Guess Wind
- Seasonal Adjusted Wind

- Adjustments to wind stress seasonal cycle explain majority of wind-driven cooling
- Cooling driven by anomalous advection

Directional decomposition of advection

- Vertical heat fluxes primarily strengthen cooling

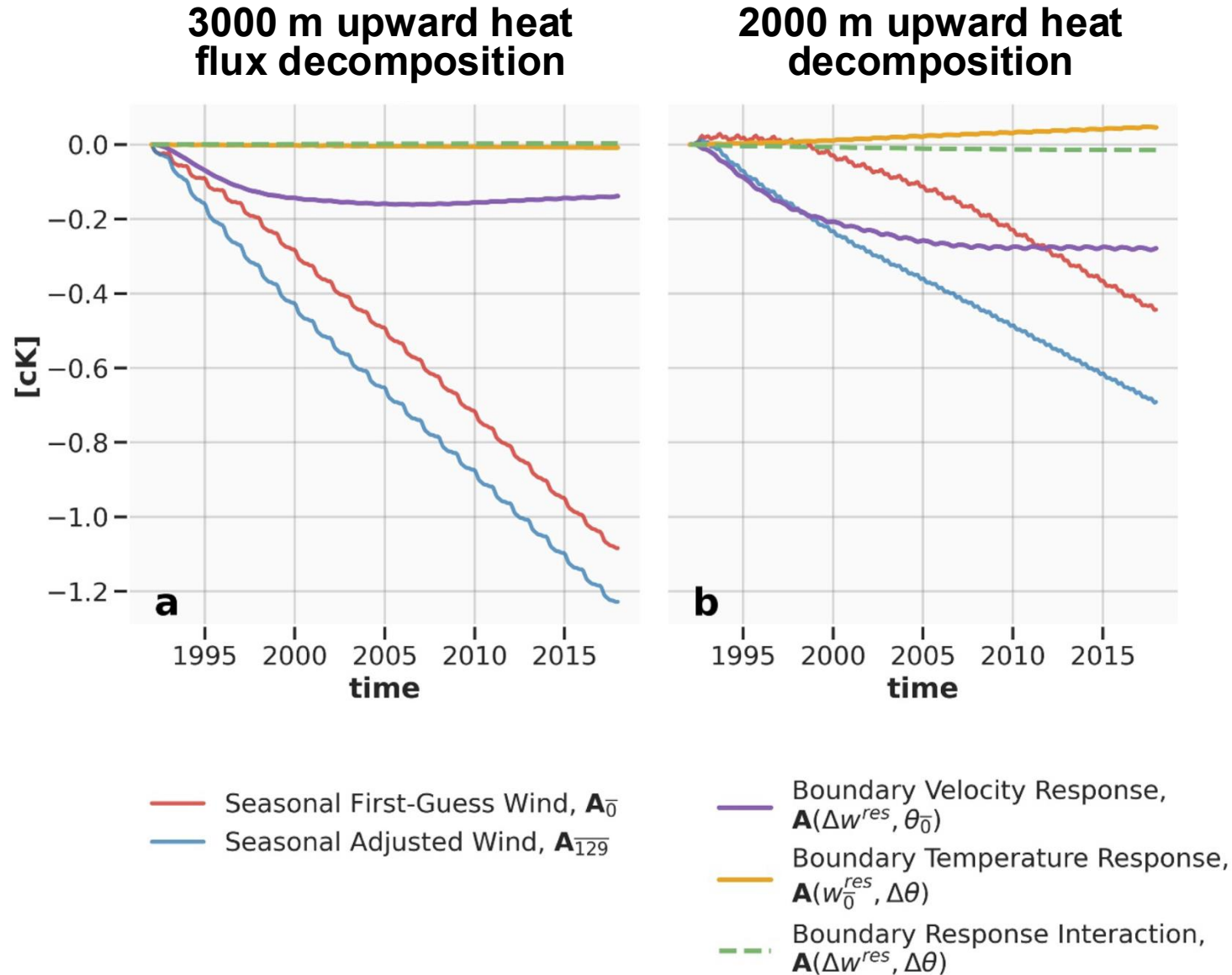


Vertical advection Reynolds decomposition

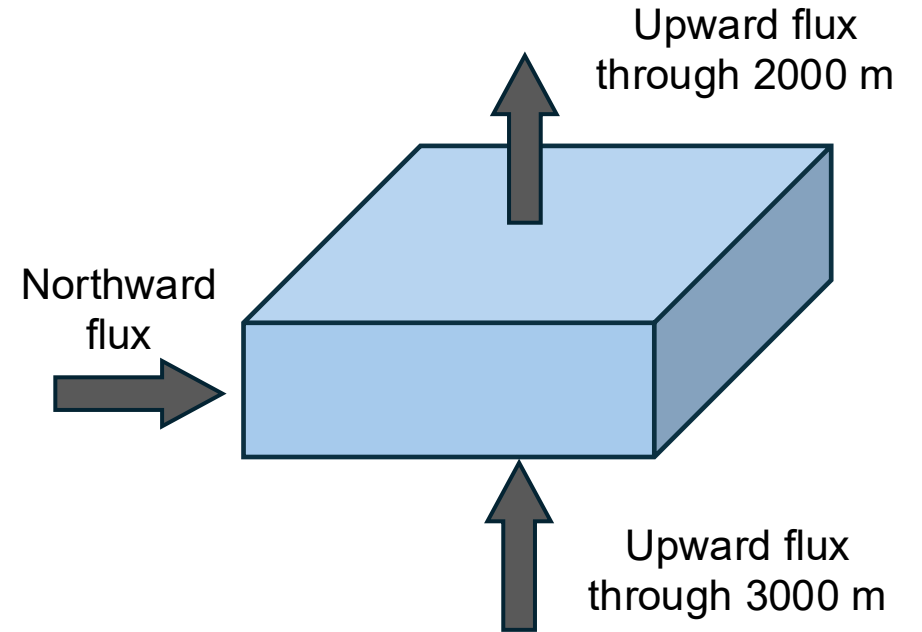
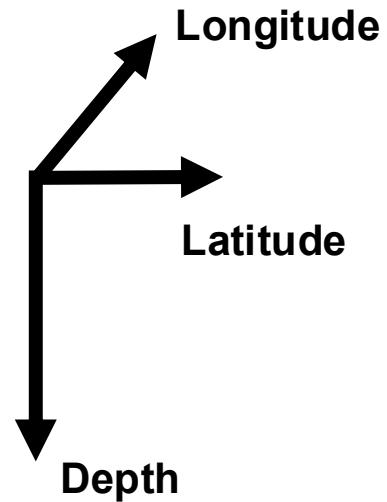
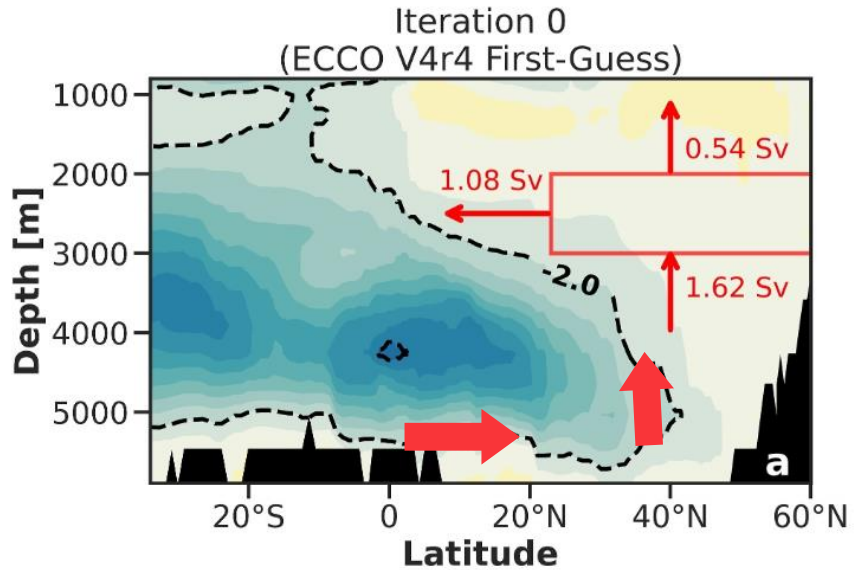
$$\Delta A = A(\Delta w, \theta_0) + A(w_0, \Delta \theta) + A(\Delta w, \Delta \theta)$$

Vertical advection Reynolds decomposition

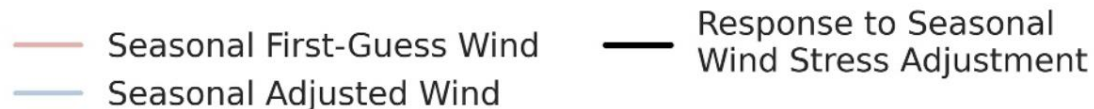
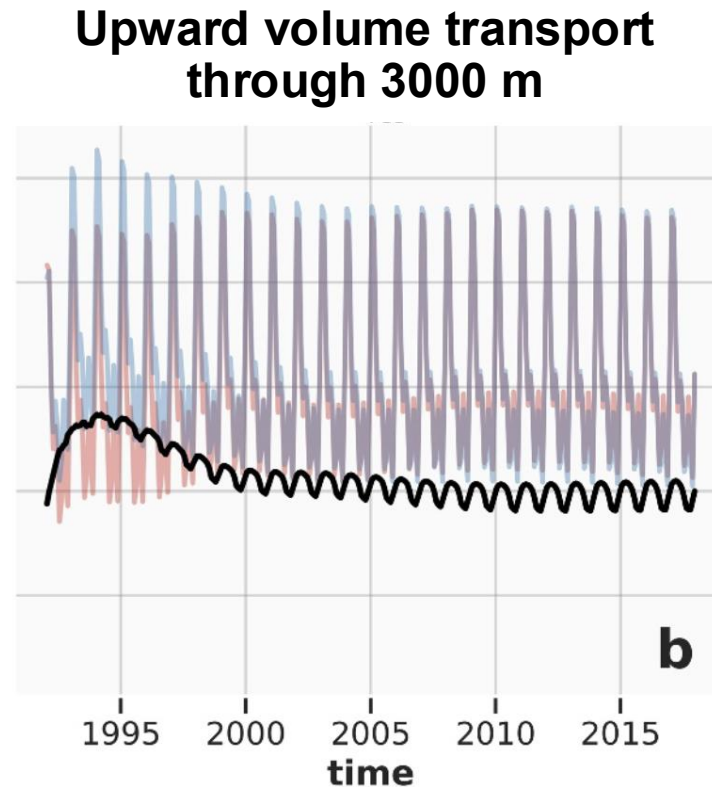
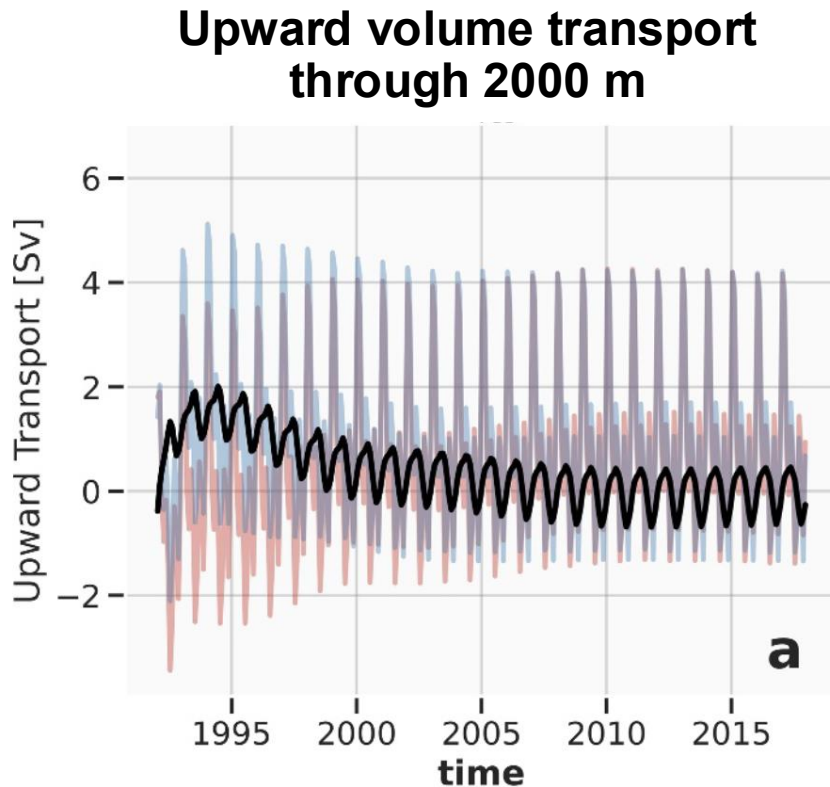
- Anomalous advection is driven by anomalous volume transport



Deep vertical transport response to winds

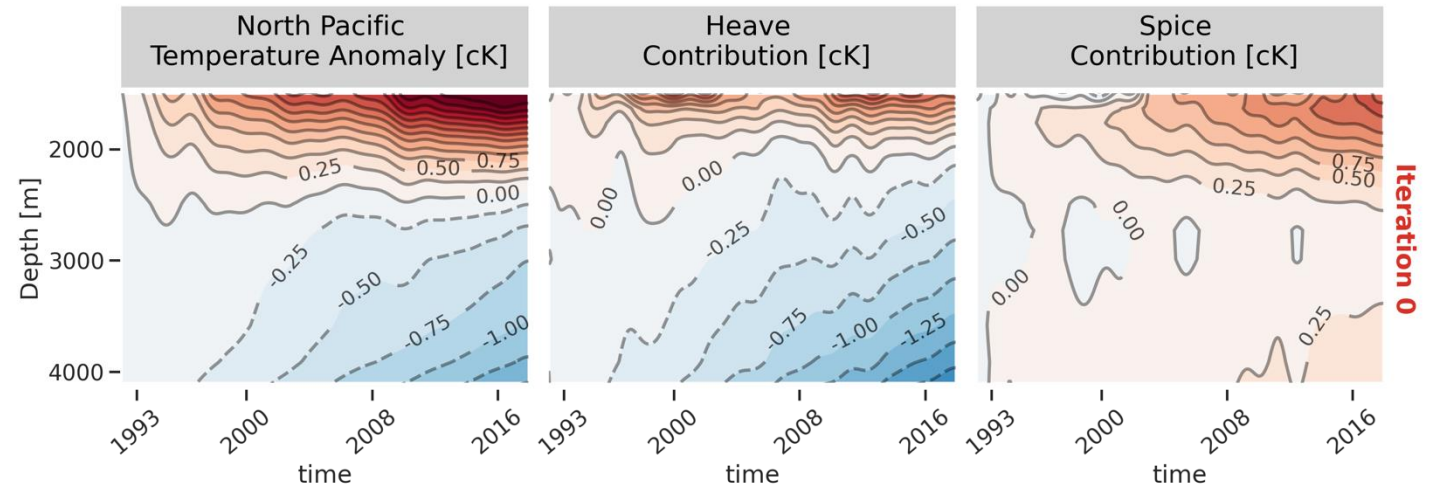


Deep vertical transport response to winds

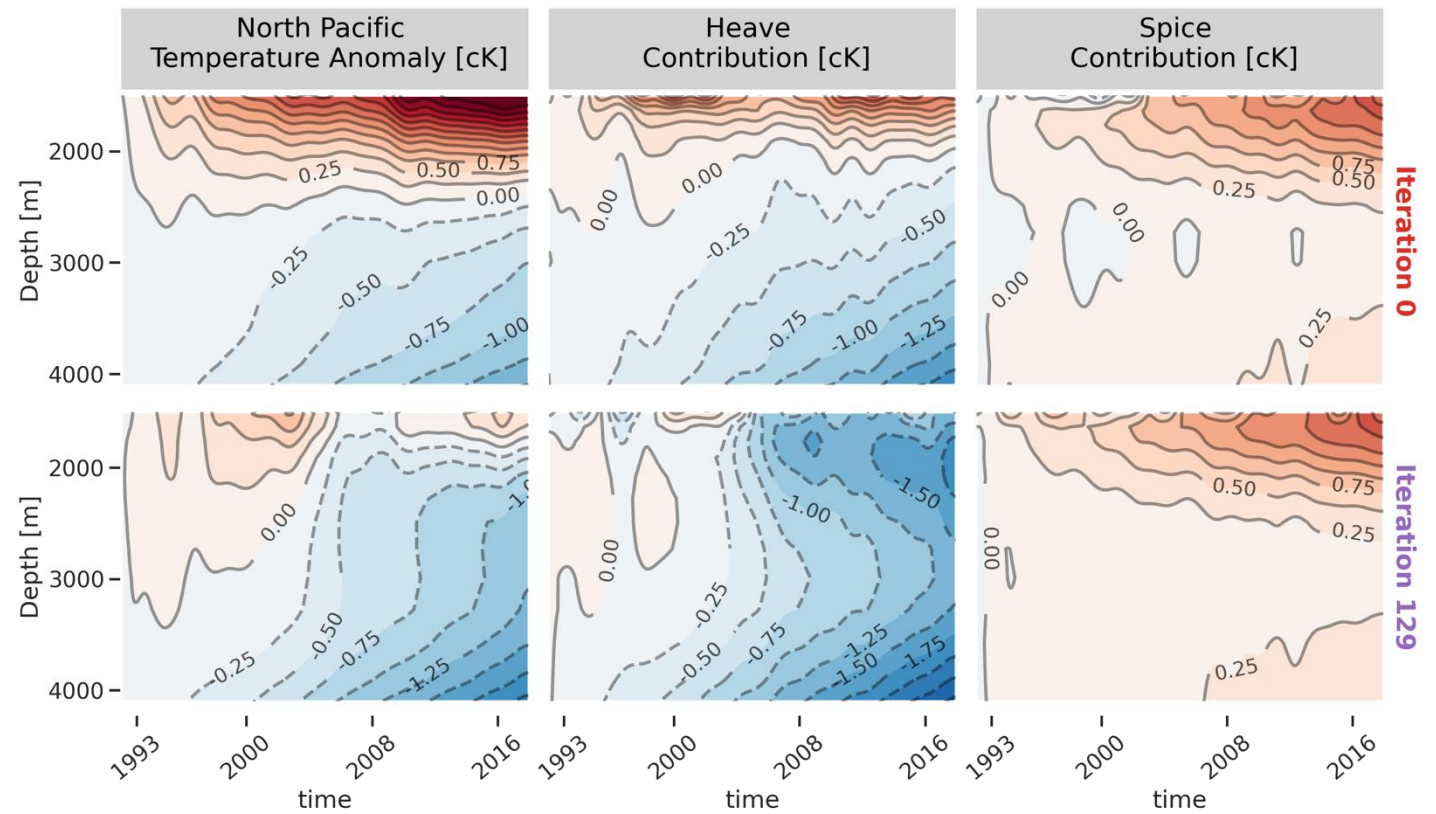


- Vertical transport anomalies are strongest during first 8 years of run
- Damped behavior consistent with Rossby wave superposition

Effects of isopycnal heave and spice



Effects of isopycnal heave and spice



Effects of isopycnal heave and spice

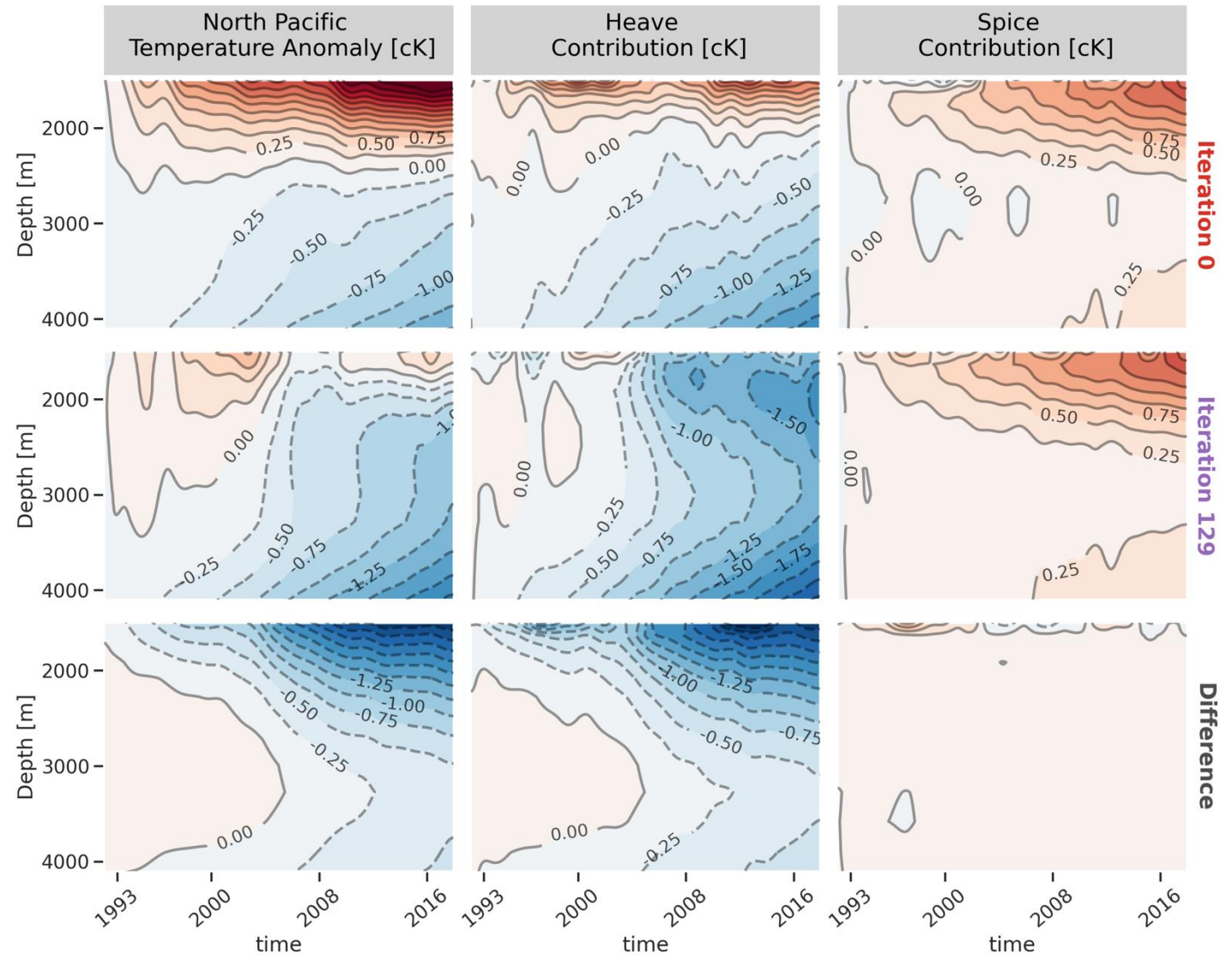
Wind-induced
vertical transport anomalies



Isopycnal upwelling



Mid-depth cooling



Summary

- ECCO V4r4 has a 5 cK/century mid-depth Pacific cooling trend resembling trends inferred from HMS *Challenger* observations
- Cooling only emerges after ECCO V4r4 is constrained to observations
- The data-constrained wind fields cause adiabatic isopycnal heave that upwells cool waters

Takeaways

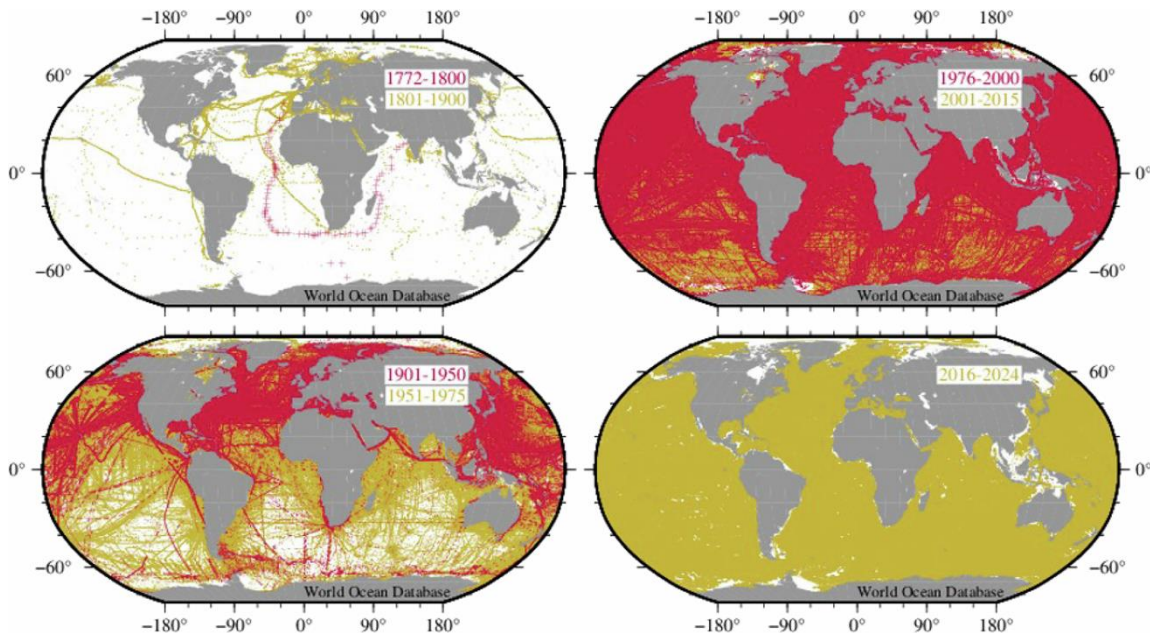
- Wind can affect deep ocean temperature trends through planetary wave adjustment
- **Adiabatic heave should be considered when interpreting deep ocean temperature records**

Where do we go from here?

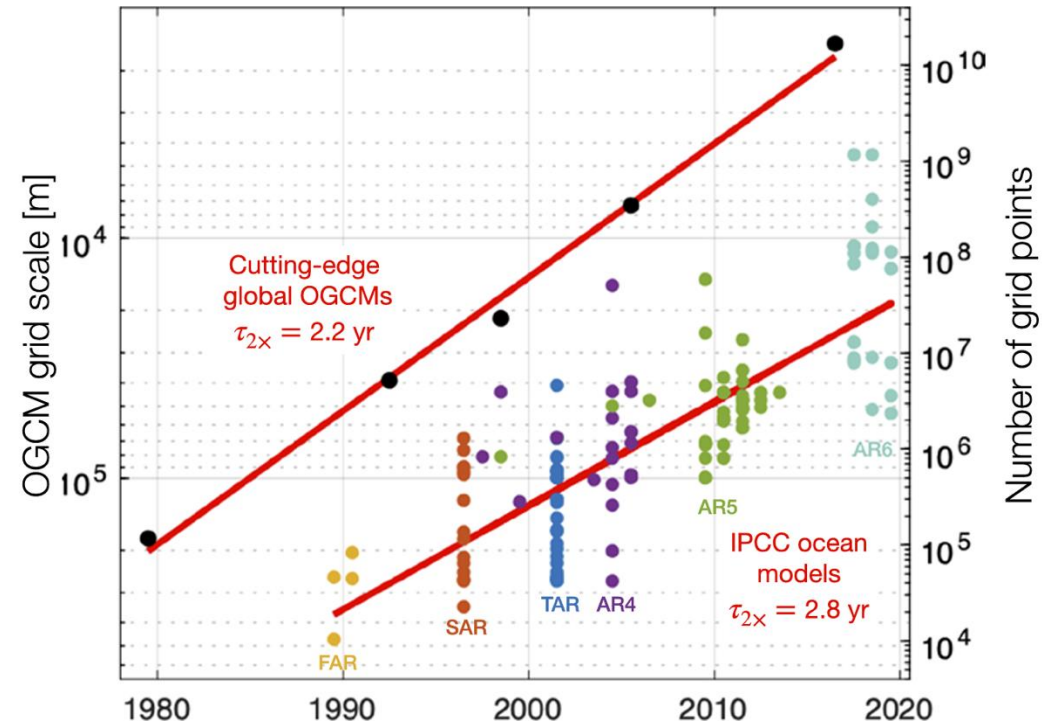
Is Computational Oceanography Coming of Age?

Thomas W. N. Haine, Renske Gelderloos, Miguel A. Jimenez-Urias, Ali H. Siddiqui, Gerard Lemson, Dimitri Medvedev, Alex Szalay, Ryan P. Abernathey, Mattia Almansi, and Christopher N. Hill

Unprecedented ocean observing *Garcia et al., 2026*



Exponential growth of OGCM resolution *Haine et al., 2021*



Thank you! 😊

The River Thames (in London)
during the Little Ice Age



The River Thames today

